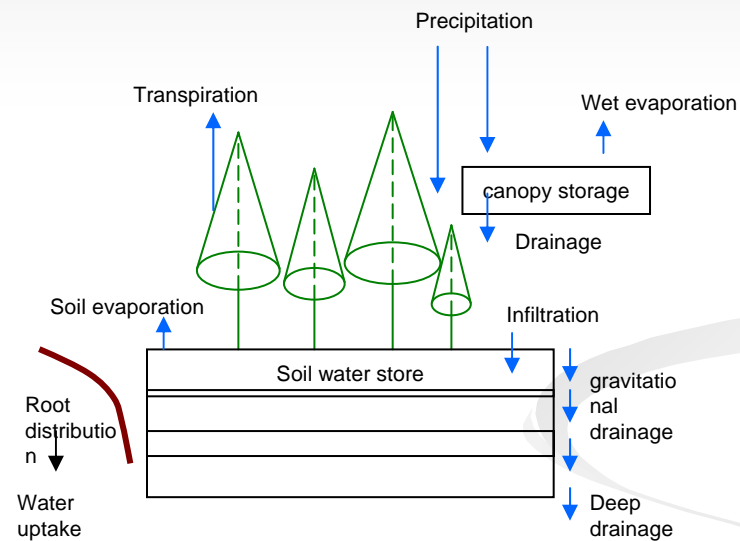


MAESPA: Development of a soil-plant-atmosphere model



Remko Duursma, Dec. 12 2008 UWS

Process-based models of canopy function

- Canopy-level gas exchange ($\text{H}_2\text{O}, \text{CO}_2$) notoriously hard to measure
 - before whole-tree chambers, eddy flux, no real way to measure
- Use of models to scale up leaf-level fluxes (which are ‘easy’ to measure) to the canopy goes back to C.T. de Wit’s “Photosynthesis of leaf canopies”, 1965.

Process-based models of canopy function

- Why complex over-parameterized models?
 - Similar predictive power can be reached with simple models
- Complex models very useful **research tools** to integrate detailed knowledge, test hypotheses, study system behaviour
 - “what-if” analyses
 - do measurements add up
 - scaling up (spatially and temporally)
 - exploring different process hypotheses

Process-based models of canopy function

- Complex models should be flexible
 - Not view it as a black box where all processes are represented in the ‘best possible way’
 - Our understanding is incomplete, models should have options to test different mechanisms (submodels)

Outline

- Development of MAESPA
 - Details, details
 - Implementation
- How and why do we use models like this?

MAESPA Components

- Scaling of leaf gas exchange (CO_2 , H_2O) to the tree canopy (**MAESTRA**)
 - Radiation extinction and leaf physiology models
- Respiration (**MAESTRA**) (leaf + woody biomass)
- Stand water balance (**SPA**)
 - Rainfall interception, infiltration and drainage, soil evaporation, water uptake
- Soil energy balance (**SPA**)
 - Used in estimating soil evaporation, not mandatory
- Not : growth, allocation (**MATE**), N cycle (**G'DAY**), snow/ice routines (**SPA**).

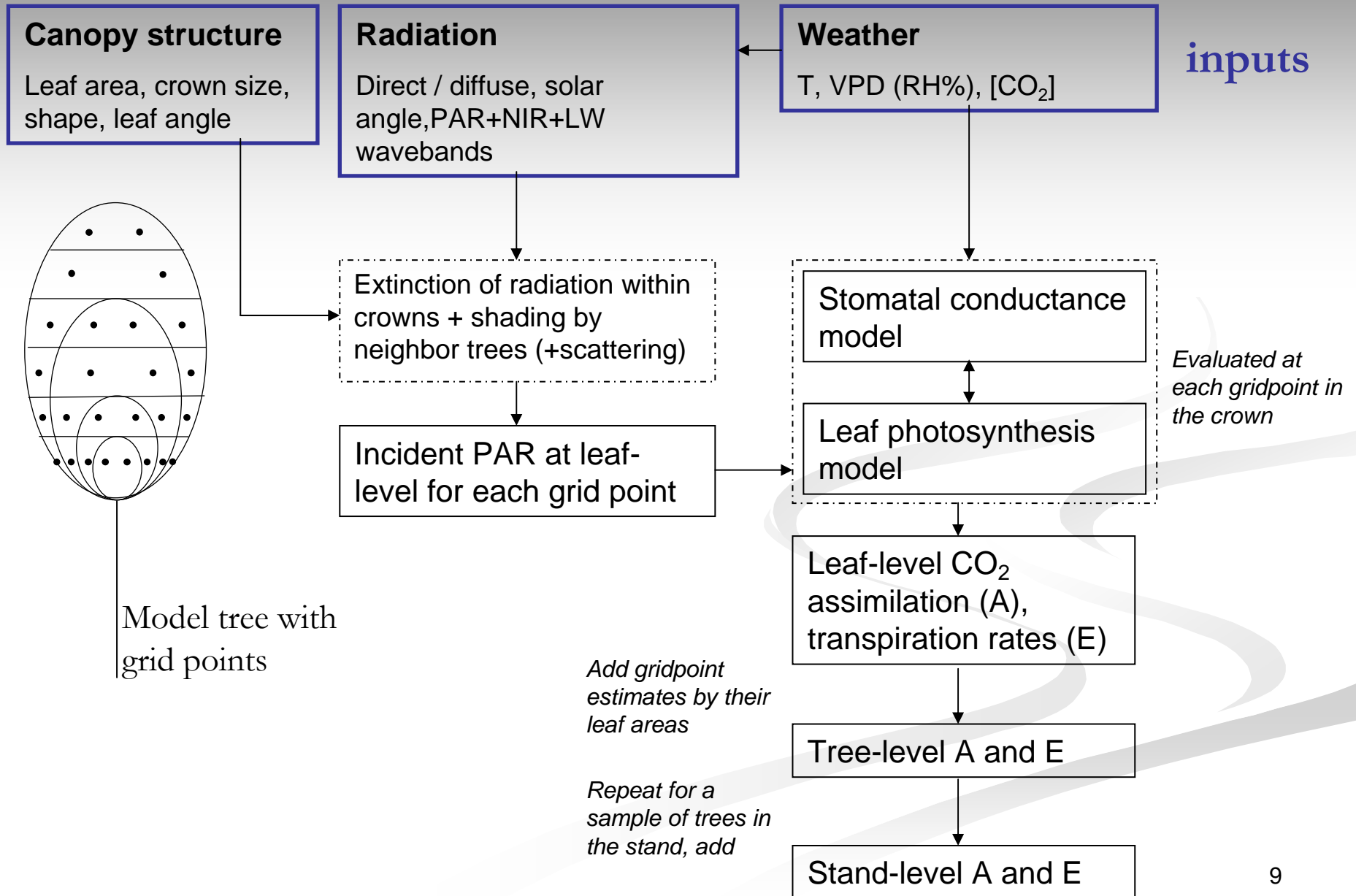
A Brief History of MAESTRA

- Norman & Jarvis (1974,1975): developed models to predict penetration of radiation in canopies, and effect of canopy structure
- Grace (1987) : important development of the radiation model, with influence from Norman and Welles.
- Wang and Jarvis (1990) : publishes MAESTRO as result of his Ph.D. research with Paul Jarvis
- Belinda Medlyn: re-organized the original Fortran code, added many options, and renamed it MAESTRA
- 50+ publications using MAESTRA

A Briefer History of SPA

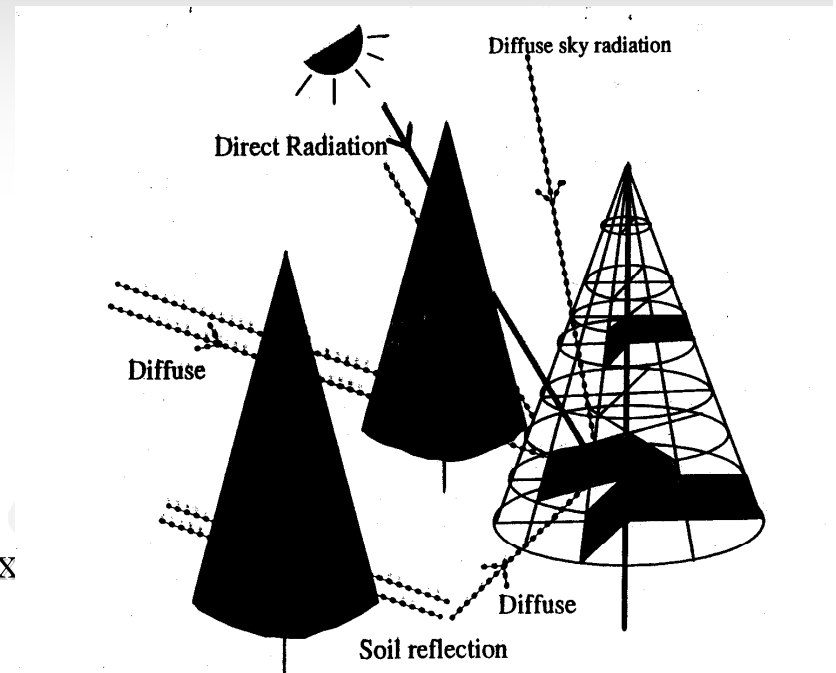
- Soil-Plant-Atmosphere model developed by Mat Williams (Williams et al. 1996, 2001a, 2001b)
- Horizontally homogenous canopy (as is usual), but a detailed coupled water and energy balance
- SPA also written in Fortran, mechanistic detail good match to MAESTRA

MAESTRA



Radiation penetration

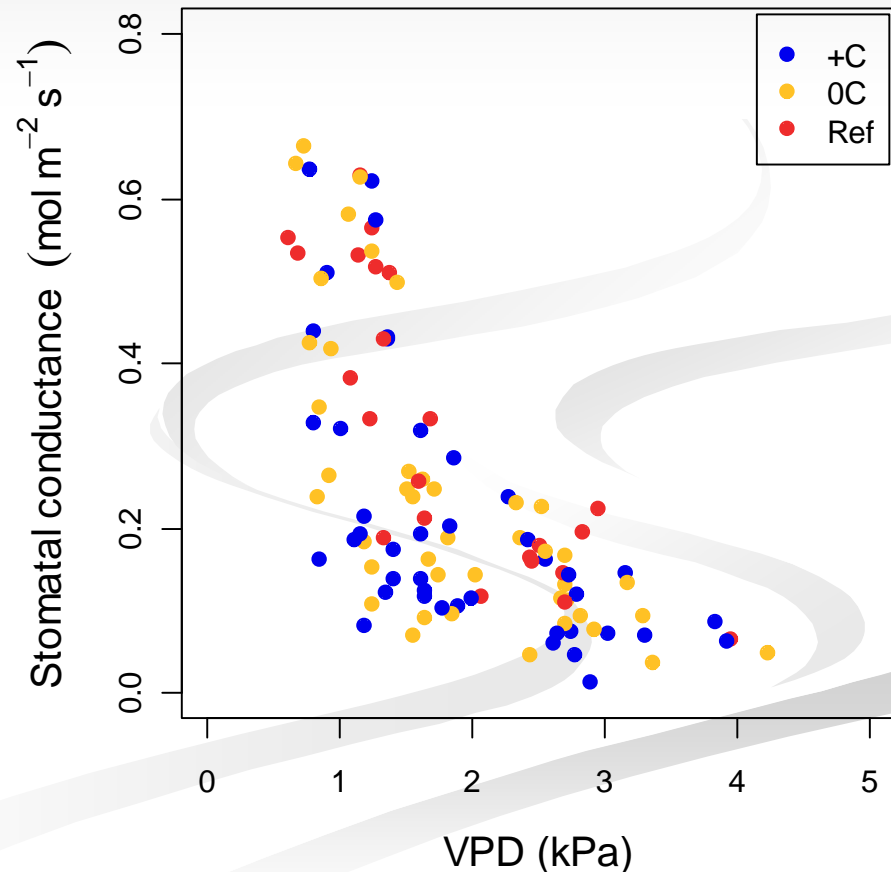
- Shading within trees, and between trees
- At each grid point, estimation of PAR, NIR, long-wave radiation
- Data needed:
 - leaf angle distribution, leaf reflectance, clumping of foliage (conifer shoots)
 - crown size (length, width), shape (ellipsoid, paraboloid, cone, cylinder, box)
 - position of neighbour trees
 - latitude, incident radiation
 - vertical and horizontal distribution of foliage in crowns (or assume evenly filled)



Stomatal conductance models (1)

- $g_s = f(\text{PAR}, \text{VPD}, \text{CO}_2, A)$
- Several options:
 - Jarvis model
 - Ball-Berry
 - Ball-Berry-Leuning
- ... other models easily added
 - optimal stomatal control

Data needed: leaf-level g_s , A ,
at varying PAR, VPD (CO_2)

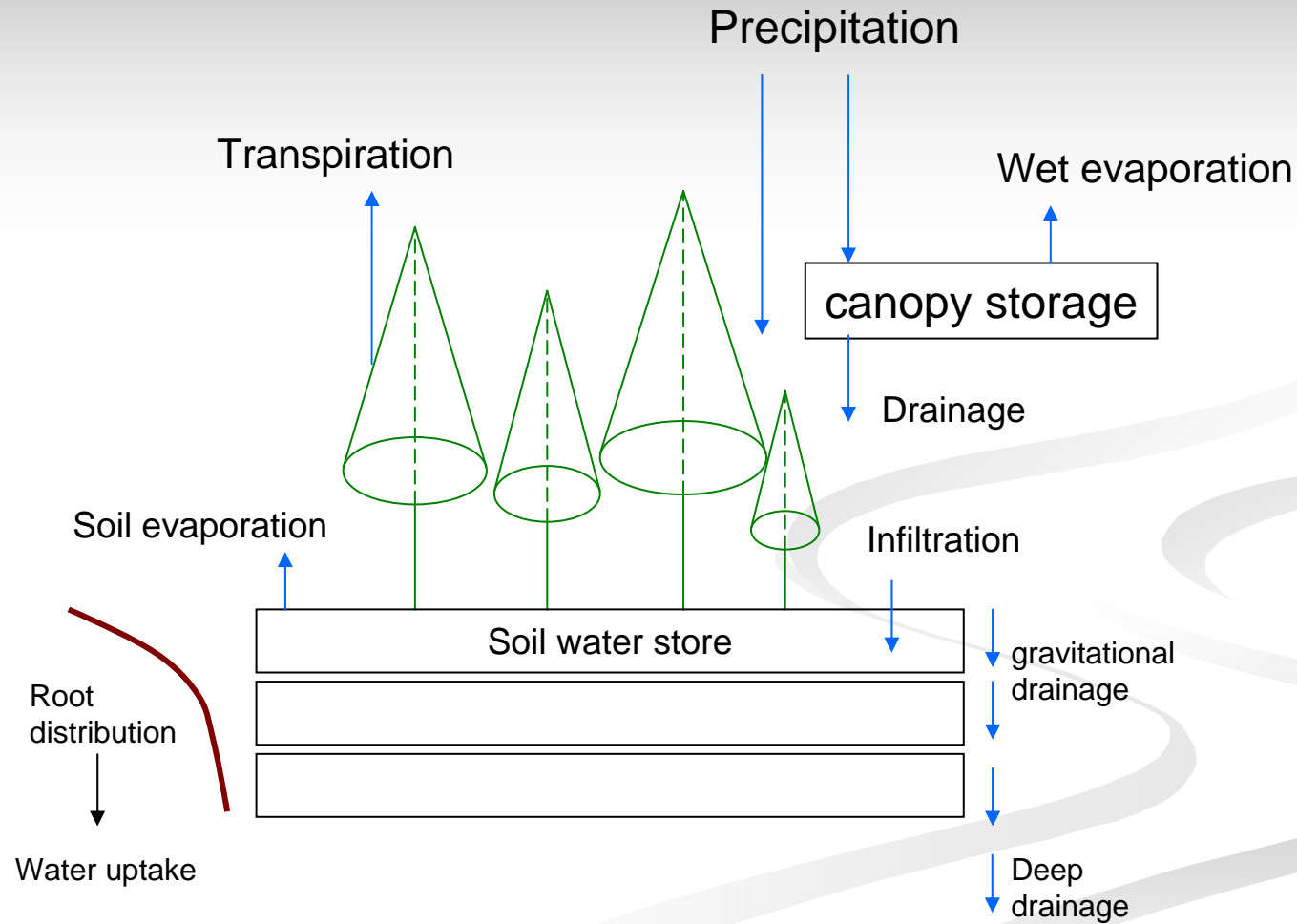


Data from Ellsworth et al., HFE

Leaf photosynthesis model

- Farquhar et al. (1980) model of photosynthesis
- Temperature dependence of V_{cmax} , J_{max} , etc.
- Quantum yield of electron transport
 - apparent quantum yield of CO_2 uptake equally as useful for parameterizing

Water balance: components borrowed from SPA



Soil evaporation

Choudhury and Monteith (1988) one-layer model

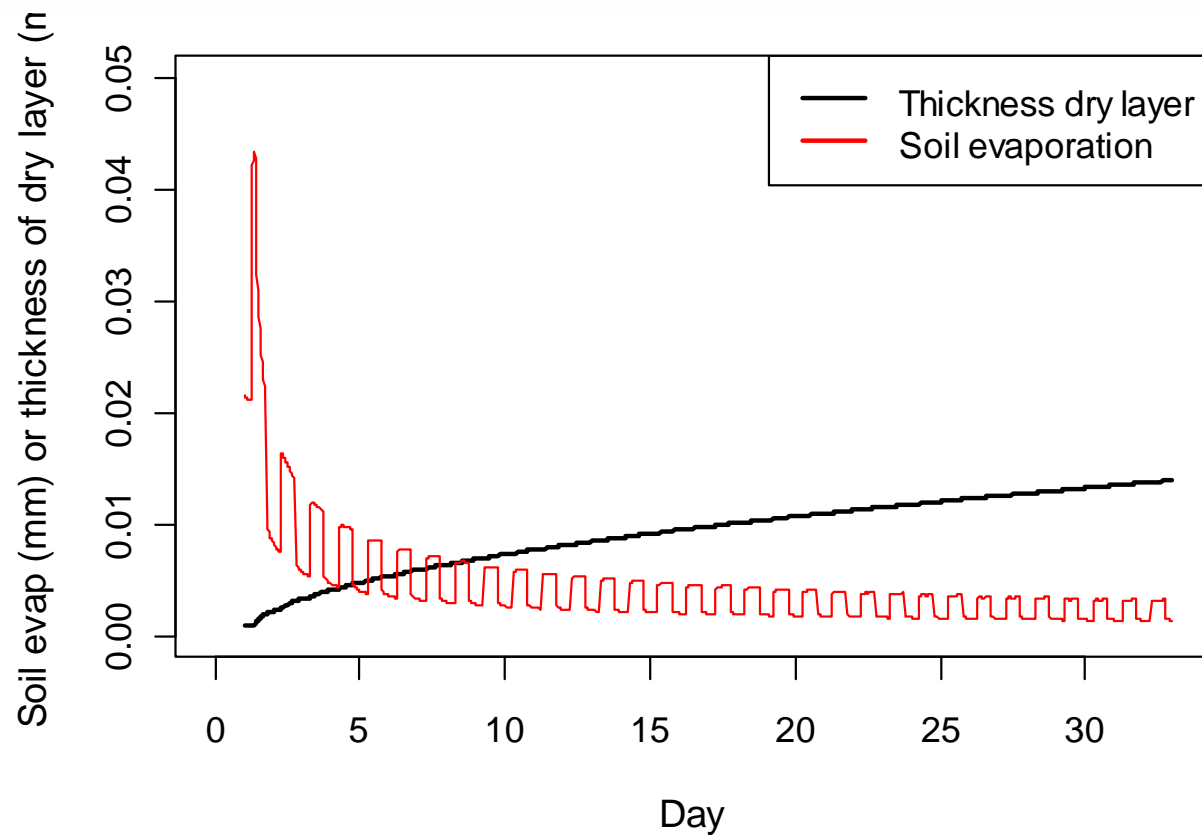
$$LE_{\text{soil}} = \alpha \frac{e_a - e_s}{r_{\text{soil}} + r_{\text{bl}}}$$

■ Where

- α – combination of (near-) constants ($\text{J m}^{-3} \text{Pa}^{-1}$)
- e_a – air vapour pressure (Pa)
- e_s – soil pore vapour pressure (= function of \mathbf{T}_{soil} and Ψ_{soil})
- r_{soil} – soil resistance (= function of **dry layer thickness**)
- r_{bl} – boundary layer resistance (s m^{-1}), function of windspeed, aerodynamic properties of canopy

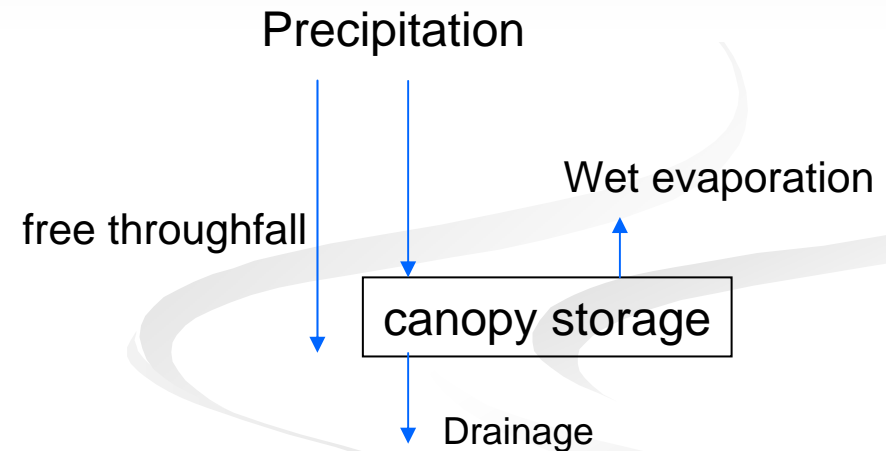
Soil evaporation (2)

- Constant weather, no rain: initial high rates of evaporation decline as dry layer increases

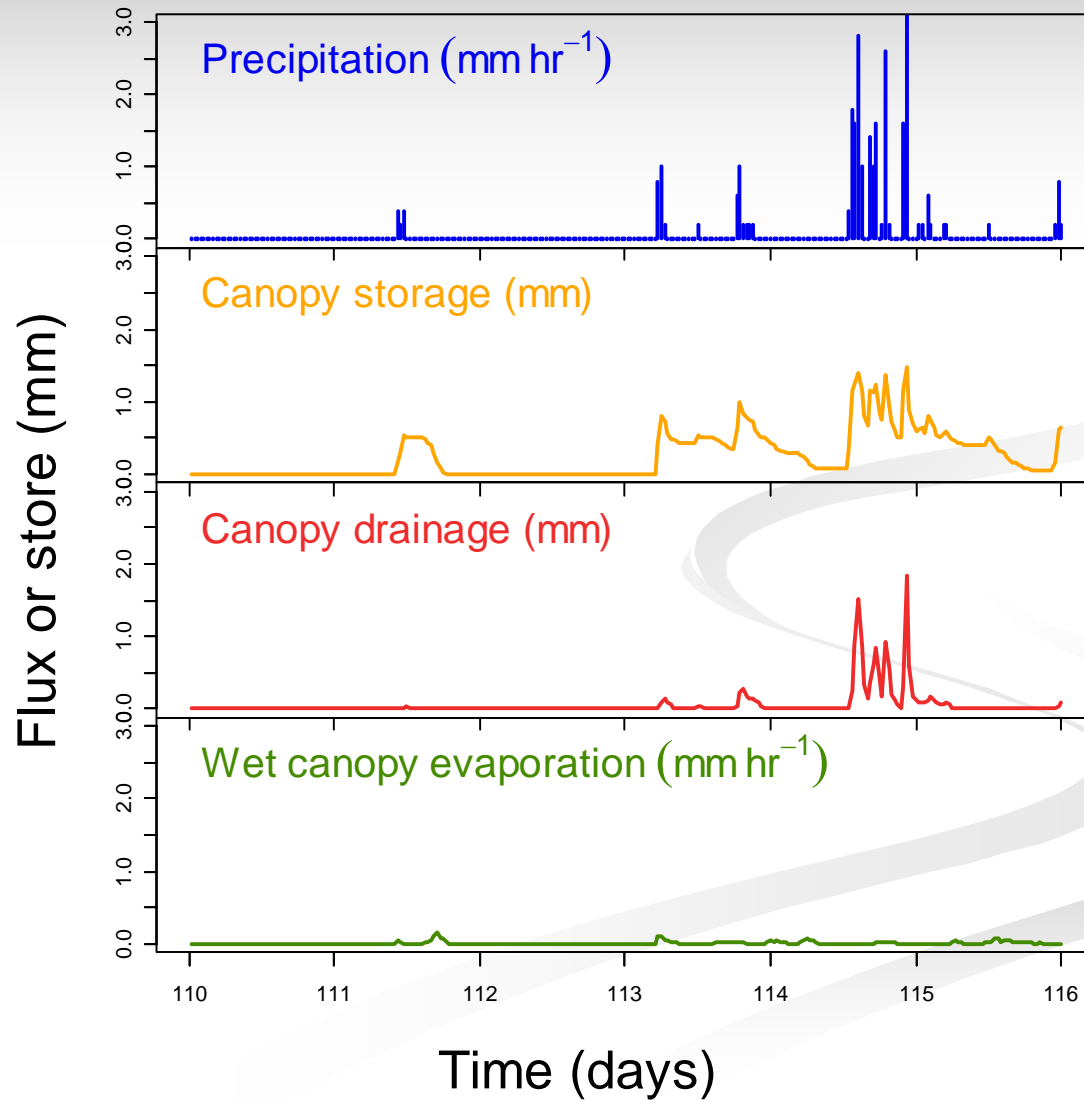


Canopy throughfall

- The classic Rutter et al. (1975) model of canopy throughfall, interception, drainage and evaporation
- Four parameters: could be derived from measurements, but often set to Rutter's defaults
- Rutter, A.J., A.J. Morton and P.C. Robins 1975. A predictive model of rainfall interception in forests. III. Generalization of the model and comparison with observations in some coniferous and hardwood stands. *Journal of Applied Ecology*. 12:367-380.



Canopy throughfall (2)

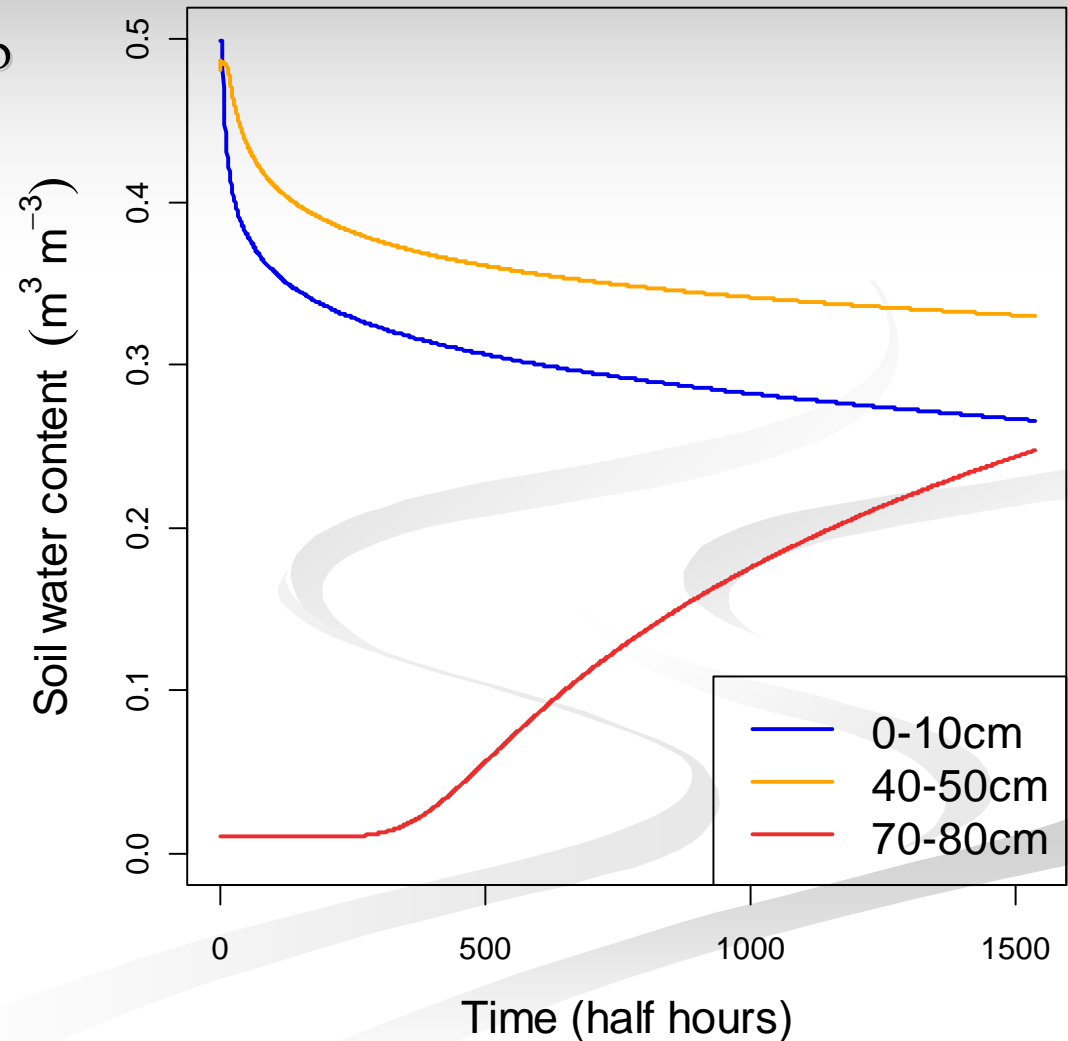


Drainage and infiltration

- Gravitational drainage is calculated from hydraulic conductivity
- Integration of the Richards' equation
 - Very standard method in soil hydrology
- No macropore flow: could be important
- Infiltration of rainfall: SPA assumed complete infiltration in top layer
- MAESPA includes option for immediate infiltration of rainfall into deeper layers (macropore idea, based on BROOK90 model)

Drainage (2)

- Example simulation: 1m deep soil, 10 layers
- Top 5 layers saturated at $t=0$
- Bottom 5 layers dry
- No transpiration

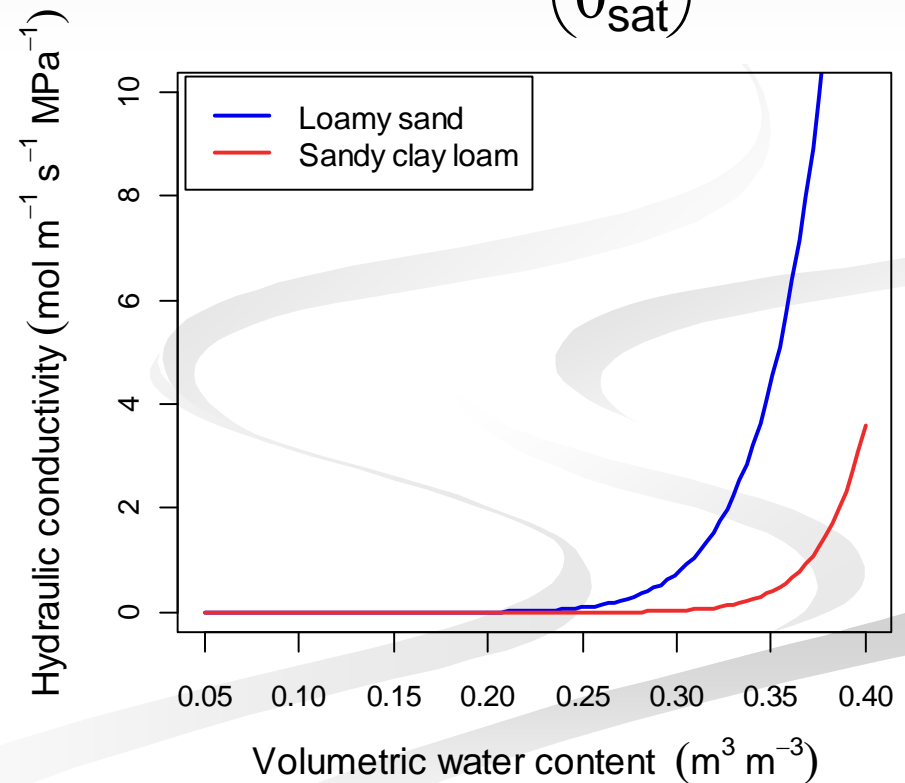
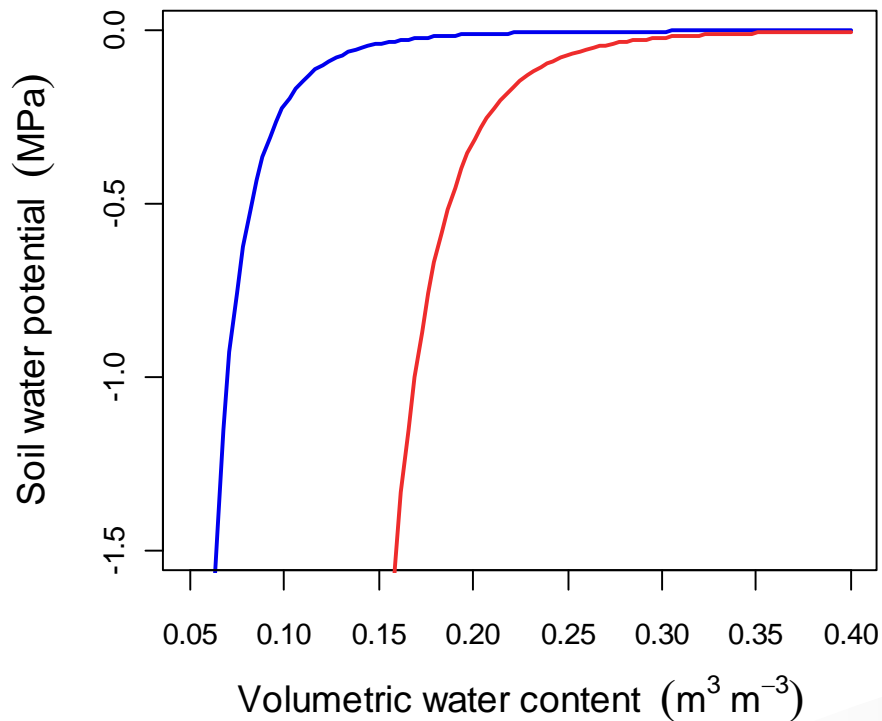


Soil water potential and hydraulic conductivity

- Campbell's (1974) coupled retention and conductivity curves

$$\Psi_{\text{soil}} = \Psi_e \left(\frac{\theta}{\theta_{\text{sat}}} \right)^{-b}$$

$$K_{\text{soil}} = K_{\text{sat}} \left(\frac{\theta}{\theta_{\text{sat}}} \right)^{2b+3}$$



Limits on leaf transpiration: one option



- Ohm's analogy to water flow, one-dimensional
- Assumption of a critical minimum leaf water potential
- Maximum transpiration rate is then:

$$E_{\max} = k_{\text{tot}} \times (\Psi_{\text{soil}} - \Psi_{\text{min}})$$

where k_{tot} conductance from soil to leaf, Ψ_{min} minimum leaf water potential, Ψ_{soil} a weighted soil water potential (by the layers)

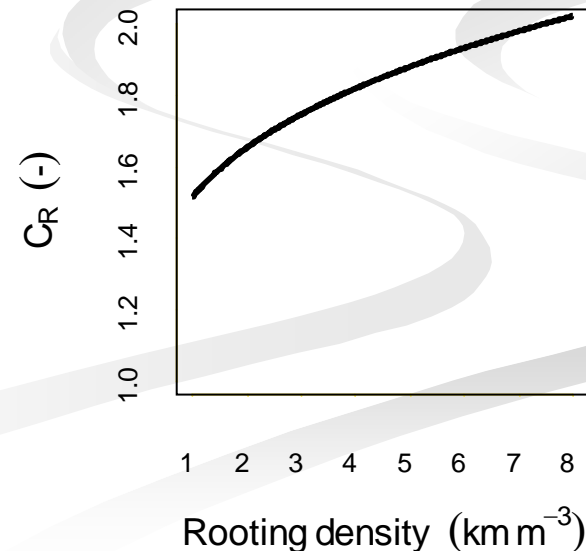
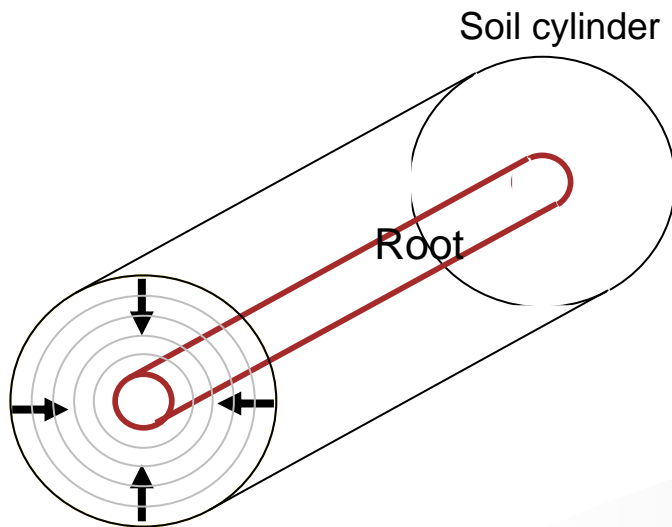
If calculated E from the stomatal conductance model exceeds E_{\max} ,
 E is set to E_{\max} (and g_s and A recalculated)

Soil to root surface conductance

- Gardner's (1960) single root model

$$k_{\text{soil}} = \frac{\text{RLI}}{\text{LAI}} \times C_R \times K_{\text{soil}}(\Psi_{\text{soil}})$$

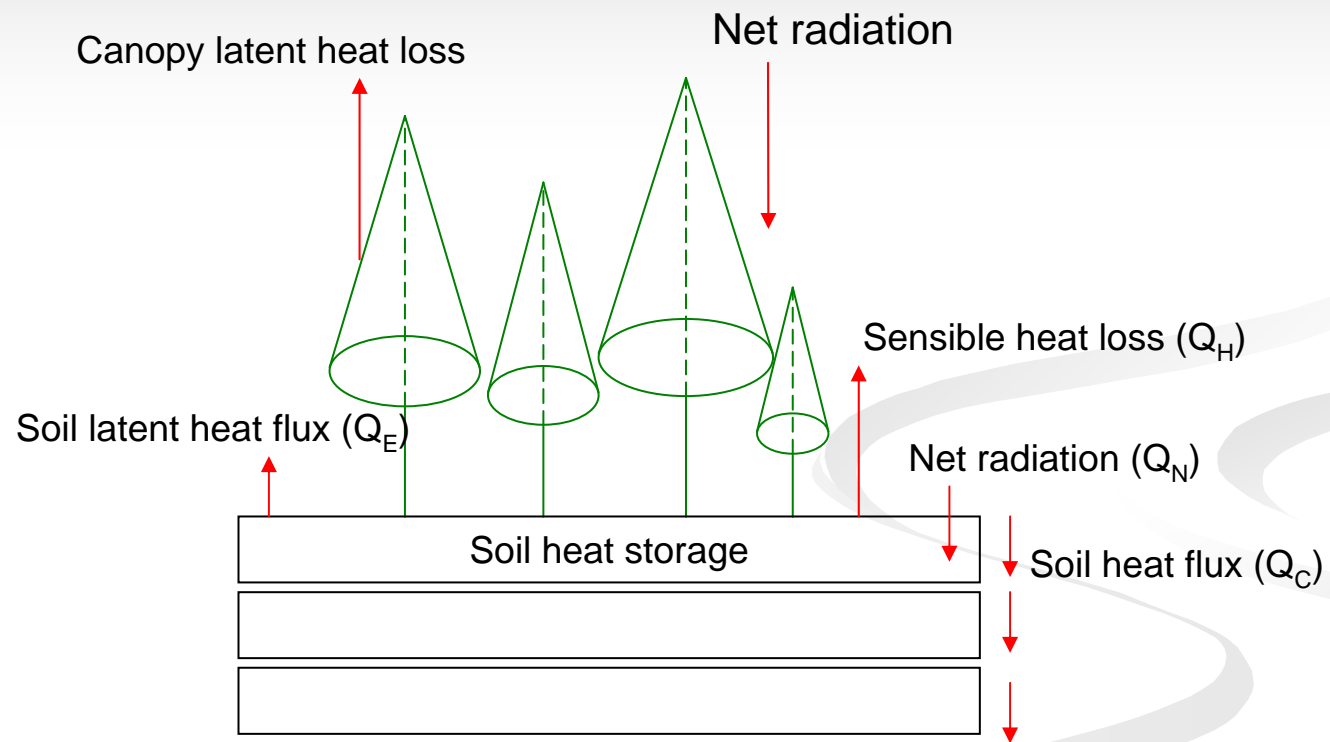
where: k_{soil} leaf-specific soil hydraulic conductance ($\text{mol m}^{-2} \text{s}^{-1} \text{MPa}^{-1}$), **RLI** root length index (m m^{-2}), **LAI** leaf area index, **K_{soil}** hydr. conductivity, **C_R** a root index function



Soil water uptake (3)

- Fraction uptake in each of the soil layers is determined from soil conductance in each layer
 - This is a SPA hypothesis, and should be tested more! Other alternatives may exist as well
- Data needed for the soil water uptake module:
 - Plant hydraulic conductance (leaf-specific) (from sapflux and drop in leaf water potential).
 - Minimum leaf water potential (MPa)
 - Soil water retention data (or soil texture at the least), saturated hydraulic conductivity
 - Rooting depth, rooting density, vertical profile

Energy balance



Energy balance (2)

- Soil surface temperature (T_s) is calculated from closing the energy balance equation:

$$R_n + Q_e + Q_h + Q_c = 0$$

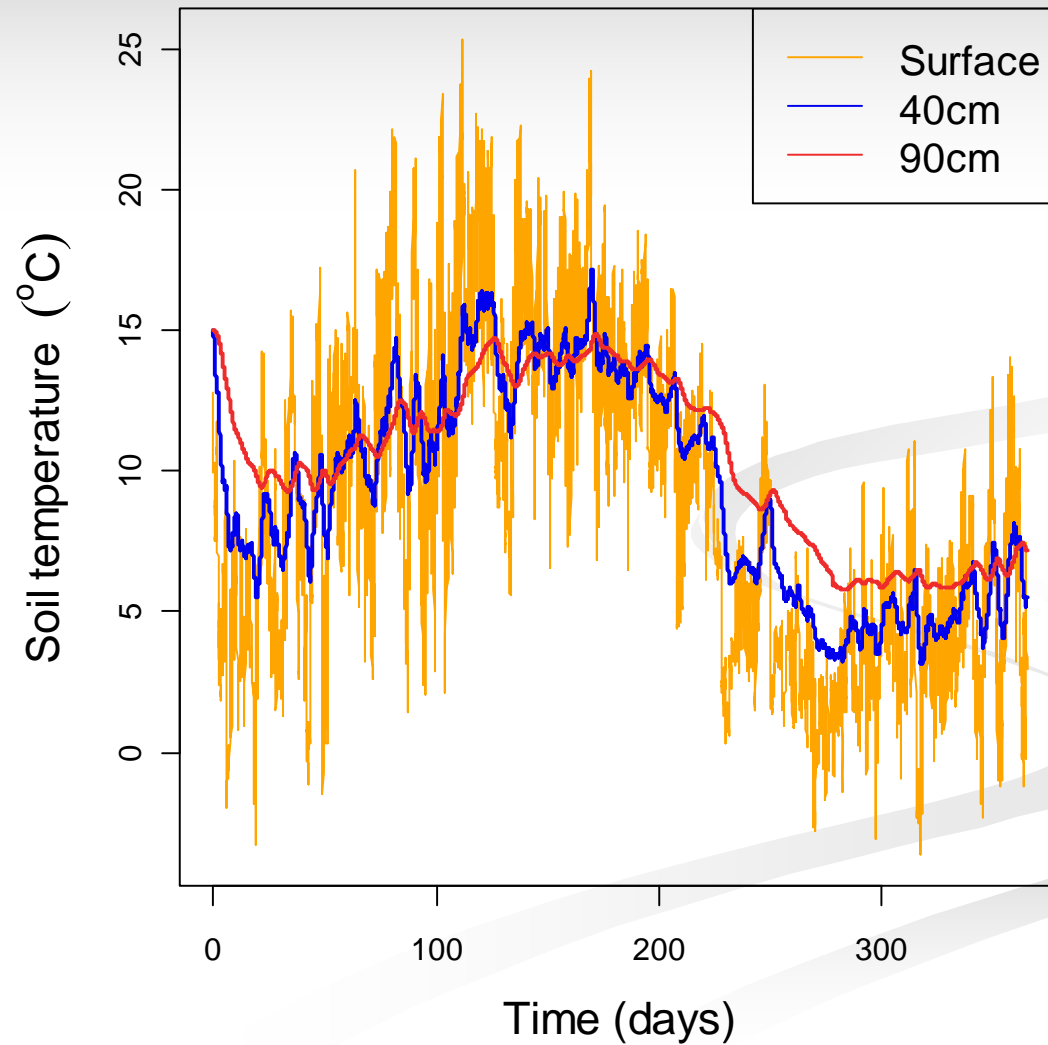
Where R_n is net radiation, Q_e latent heat loss (soil evaporation), Q_h soil heat flux, Q_c sensible heat loss, all in $W m^{-2}$.

- All heat fluxes depend on T_s , so it is possible to solve the energy balance equation for T_s
- Soil evaporation is then calculated from this surface temperature

Soil heat flux and temperature profile

- Flux of heat in the soil depends on soil thermal conductivity
 - Function of water content, porosity, organic matter content (Lu et al. 2007)
- Litter layer is 100% organic matter, has very low conductivity
- Given the thermal conductivity for each layer, and their temperatures, we can calculate the flux of heat between layers
- This gives the soil temperature profile
- Solution of the so-called Fourier heat transport equation, standard method
- Lu, S., T. Ren, Y. Gong and R. Horton 2007. An Improved Model for Predicting Soil Thermal Conductivity from Water Content at Room Temperature. *Soil Sci Soc Am J.* 71:8-14.

Soil temperature profile example

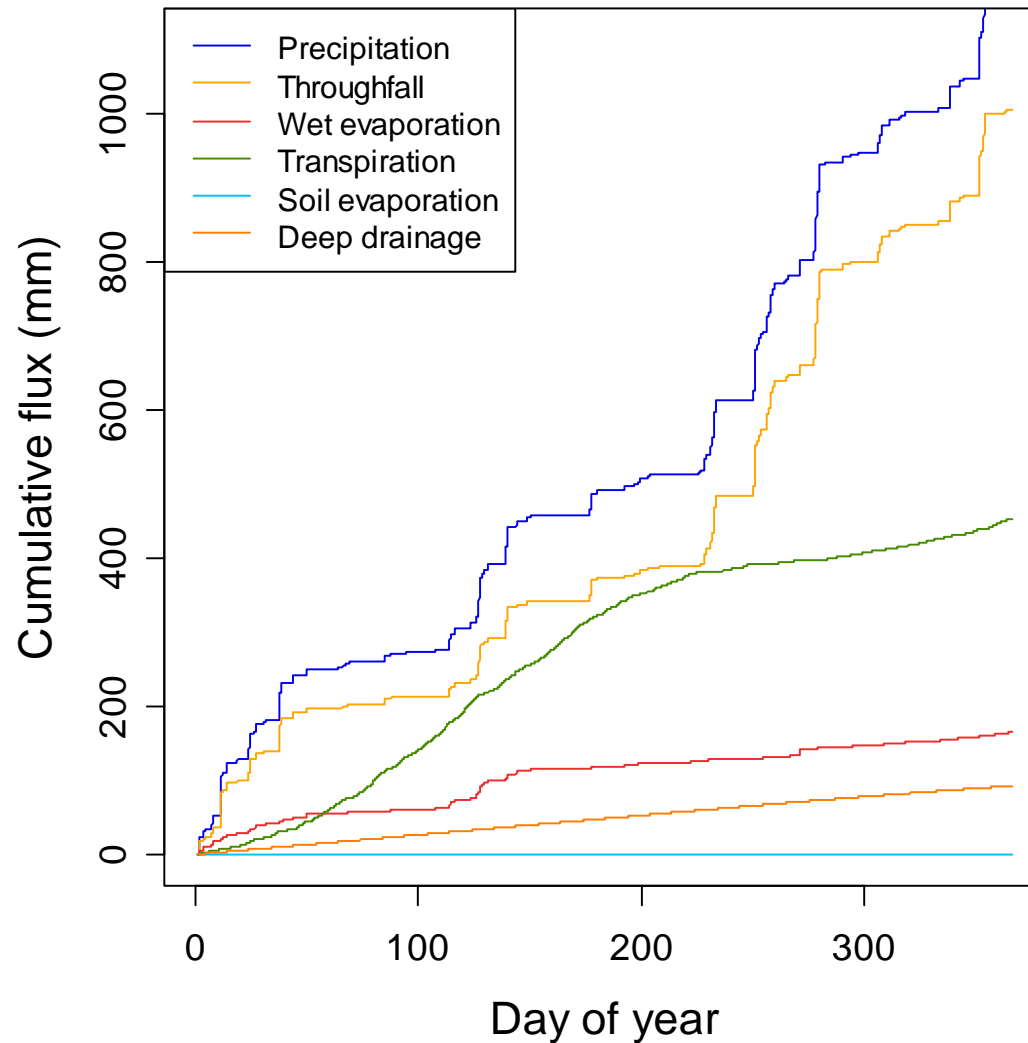


Implementation and user interface

- MAESPA is written in Fortran (as are MAESTRA, SPA)
- SPA code heavily re-organized, style and functionality matches MAESTRA
- Input text files, input error checking
 - One file for water balance parameters (`watpars.dat`)
- Output files:
 - (Half-)Hourly water balance file (`watbal.dat`)
 - Soil temperature profile (hourly) (`watsoilt.dat`)
 - Relative water uptake profile (hourly) (`watupt.dat`)
 - Water content by layer (hourly) (`watlay.dat`)

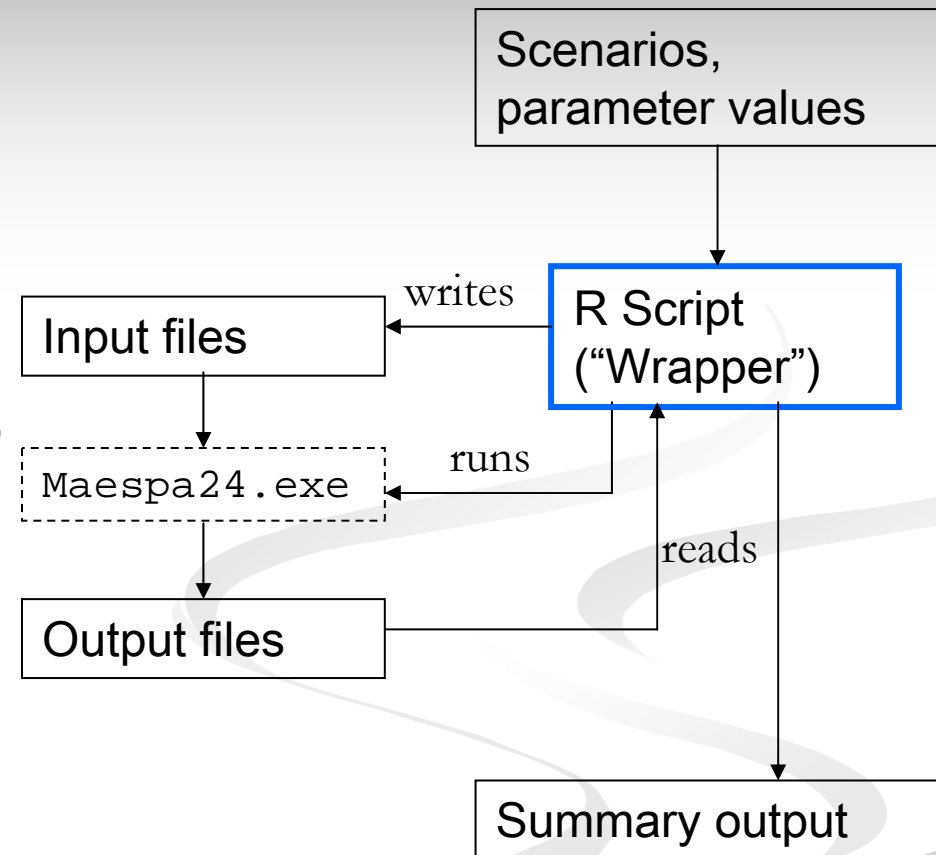
Example output (watbal.dat)

- Tumberumba flux site, *Eucalyptus delegatensis*, LAI = 1.5 m² m⁻²



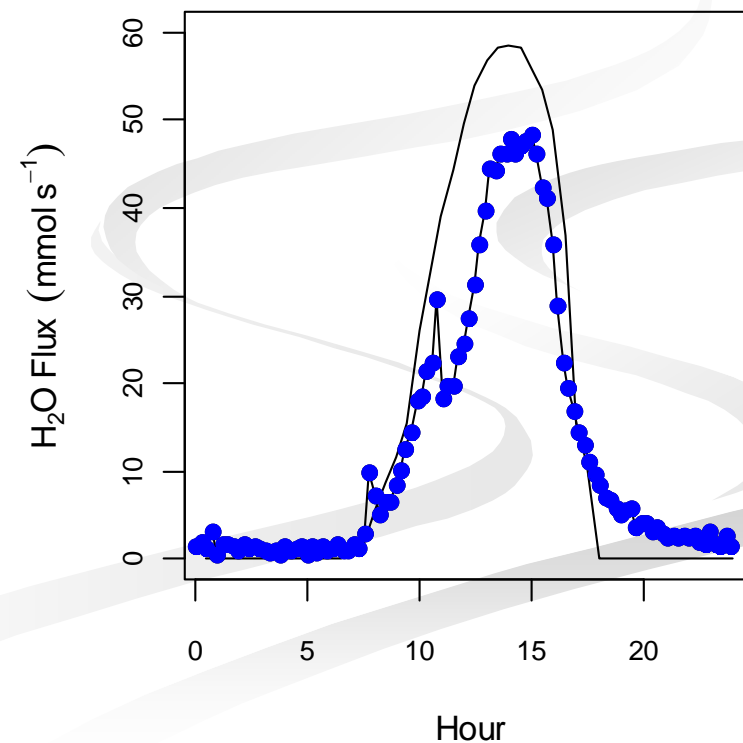
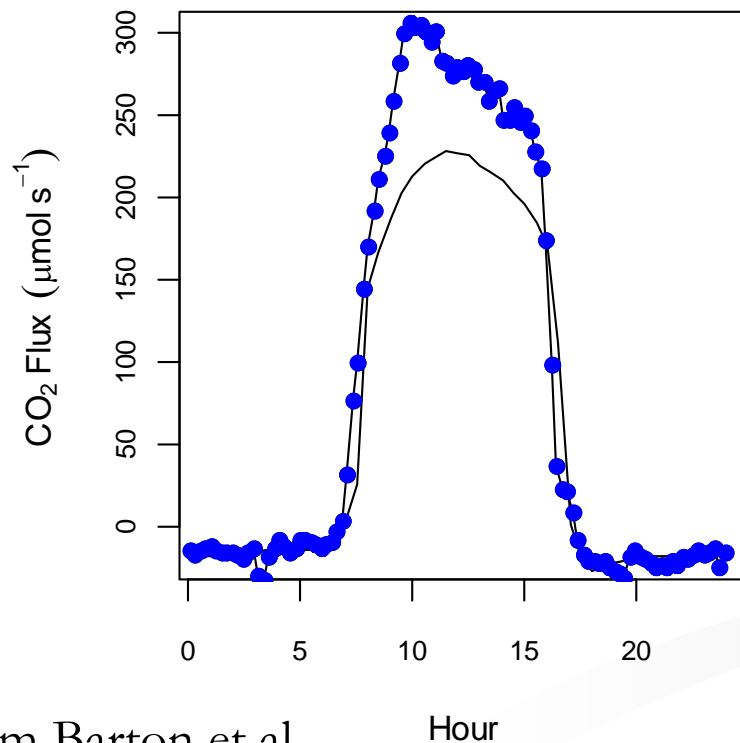
Batch Utility

- A collection of functions written in R for multiple simulations of MAESPA/MAESTRA
- Can be used for other models that use namelists in input files (Fortran)
- Available as an R package
- Also available a number of functions that graph MAESPA output



HFE

- Whole-tree fluxes of CO_2 and H_2O ideal for model testing
- Very preliminary runs: no competitive shading, stomatal conductance model not properly calibrated.
- One Saturday's data in April 2008 – total leaf areas estimated in April



Data from Barton et al.

Hour

Hour

Potential model applications at the HFE

- Where is the water coming from?
 - Do fluxes of water add up?
 - Depth of water uptake?
 - Can we predict effects of drought treatment?
- CO₂ effect: contributions of LA, LUE, g_s, etc.
- Testing mechanisms: poor understanding of drought
- Are leaf-level measurements consistent with whole-tree fluxes?
- Impacts of reduced g_s on soil water balance, sensitivity to drought : do model predictions match tree fluxes?

Other applications

- Testing at other sites (eddy-flux, sap-flux sites)
- Development of response surfaces to aid development and parameterization of simple models (e.g. MATE, MATEY)
 - Scaling up from the HFE to...
- Sensitivity analyses:
 - important parameters?
 - What-if? (e.g. “What if there would be no downregulation?”)
- Strength of MAESTRA: 3D canopy structure
 - Effect on energy and water balance?
 - Do simple models need to be adjusted?